Passive correlation based imaging

Chrysoula Tsogka

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University of Crete and IACM-FORTH

Support: ERC-StG: 239959



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- Estimating velocity change in a medium
- 8 Seasonal Variations
- Experiments on Real data

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 - K.G. Sabra, P. Gerstoft, P. Roux, W.A. Kuperman and M.C. Fehler, "Extracting time domain Green's function estimates from ambient seismic noise", Geophys. Res. Lett. 32, L03310, 2005.

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 - New developments on imaging and monitoring with seismic noise. C. R. Geosciences, 343 (8-9), 2011

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Image: A math a math

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- Applications :
 - Structural Health Monitoring
 - - E. Larose, O.I. Lobkis, and R. L. Weaver "Passive correlation imaging of a buried scatterer", J. Acoust. Soc. Am., 119 (6), 2006.

K. Sabra et al, "Structural health monitoring by extraction of coherent guided waves from diffuse fields", J. Acoust. Soc. Am. 123 (1), 2008.

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 - Exploration geophysics (auxiliary array imaging)
 - - A. Bakulin and R. Calvert, The virtual source method : Theory and case study, Geophysics, 71(2006), pp. SI139-SI150
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 - Lorent Brenguier, Daniel Clarke, Yosuke Aoki, Nikolai M. Shapiro, Michel Campillo, Vale rie Ferrazzini, "Monitoring volcanoes using seismic noise correlations", C. R. Geoscience 343 (2011) 633–638.
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- \bullet We consider a domain Ω containing a reflector ${\cal O}$
- and $u(t, \vec{\mathbf{x}})$ solution of the acoustic wave equation on $\Omega \setminus \mathcal{O}$:

$$\frac{1}{c(\vec{\mathbf{x}})^2} \frac{\partial^2 u(t, \vec{\mathbf{x}})}{\partial t^2} - \Delta u(t, \vec{\mathbf{x}}) = n(t, \vec{\mathbf{x}}).$$

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• Propagation speed $c(\vec{\mathbf{x}})$ is given

$$\begin{split} c(\vec{\mathbf{x}})^2 &= c_0^2 \\ \frac{1}{c(\vec{\mathbf{x}})^2} &= \frac{1}{c_0^2} \left[1 + \sigma \mu(\frac{\vec{\mathbf{x}}}{\ell}) \right] \end{split}$$

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- $\bullet~\sigma$ controls the strength of the fluctuations,
- $n(t, \vec{\mathbf{x}})$ models noise sources.

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Time domain

• The Green's function $G(t,\vec{\mathbf{x}},\vec{\mathbf{y}})$ is the solution of

$$\frac{1}{c(\vec{\mathbf{x}})^2} \frac{\partial^2 G(t, \vec{\mathbf{x}}, \vec{\mathbf{y}})}{\partial t^2} - \Delta G(t, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = \delta(t)\delta(\vec{\mathbf{x}} - \vec{\mathbf{y}})$$

• Assuming $G(t, \vec{x}, \vec{y}) = 0$ for t < 0 we obtain in a homogeneous medium $c(\vec{x}) = c_0$, the well-known expression

$$G(t, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = \frac{1}{4\pi |\vec{\mathbf{x}} - \vec{\mathbf{y}}|} \delta(t - \frac{|\vec{\mathbf{x}} - \vec{\mathbf{y}}|}{c_0})$$

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The Green's function

Frequency domain

• In the frequency domain, the time-harmonic Green's function

$$\widehat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = \int G(t, \vec{\mathbf{x}}, \vec{\mathbf{y}}) e^{i\omega t} dt$$

is the solution of

$$\frac{\omega^2}{c(\vec{\mathbf{x}})^2}\widehat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) + \Delta\widehat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = -\delta(\vec{\mathbf{x}} - \vec{\mathbf{y}})$$

with the Sommerfeld radiation condition (assuming $c(\vec{\mathbf{x}}) = c_0$ at infinity)

$$\lim_{|\vec{\mathbf{x}}|\to\infty} |\vec{\mathbf{x}}| \left(\frac{\vec{\mathbf{x}}}{|\vec{\mathbf{x}}|} \nabla_{\vec{\mathbf{x}}} - \frac{i\omega}{c_0} \right) \widehat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = 0$$

• In a homogeneous medium we obtain

$$\widehat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = \frac{1}{4\pi |\vec{\mathbf{x}} - \vec{\mathbf{y}}|} e^{i\frac{\omega}{c_0}|\vec{\mathbf{x}} - \vec{\mathbf{y}}|}$$

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wave equation solution

 ${\ensuremath{\, \bullet }}$ The solution of the wave equation for a source $n(t,\vec{\mathbf{x}})$ is

$$u(t, \vec{\mathbf{x}}) = \int \int G(t - s, \vec{\mathbf{x}}, \vec{\mathbf{y}}) n(s, \vec{\mathbf{y}}) d\vec{\mathbf{y}} ds$$

and in the frequency domain

$$\widehat{u}(\omega, \vec{\mathbf{x}}) = \int \widehat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) \widehat{n}(\omega, \vec{\mathbf{y}}) d\vec{\mathbf{y}}$$

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The Kirchhoff Helmholtz identity



Assume that the medium is homogeneous outside B(0, D), then

$$\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) - \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)} = \frac{2i\omega}{c_0} \int_{\partial B(0,L)} \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) ds(\vec{\mathbf{y}})$$

for $L \gg D$ and $\forall \vec{x}_1, \vec{x}_1 \in B(0, D)$ (the medium can be heterogeneous in B(0, D)).

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The Kirchhoff Helmholtz identity Proof

• Let
$$G(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1)$$
 and $G(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2)$ solutions of

$$\frac{\omega^2}{c(\vec{\mathbf{x}})^2}\overline{\widehat{G}(\omega,\vec{\mathbf{y}},\vec{\mathbf{x}}_1)} + \Delta\overline{\widehat{G}(\omega,\vec{\mathbf{y}},\vec{\mathbf{x}}_1)} = -\delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_1) \quad (1)$$
$$\frac{\omega^2}{c(\vec{\mathbf{x}})^2}\widehat{G}(\omega,\vec{\mathbf{y}},\vec{\mathbf{x}}_2) + \Delta\widehat{G}(\omega,\vec{\mathbf{y}},\vec{\mathbf{x}}_2) = -\delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_2) \quad (2)$$

• Multiply (2) by $\overline{\widehat{G}(\omega, \mathbf{y}, \mathbf{x}_1)}$ and (1) by $\widehat{G}(\omega, \mathbf{y}, \mathbf{x}_2)$ and subtract $\overline{\widehat{G}(\omega, \mathbf{y}, \mathbf{x}_1)} \times (2) - \widehat{G}(\omega, \mathbf{y}, \mathbf{x}_2) \times (1)$:

$$\begin{split} \nabla \cdot \left(\overline{\widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1)} \nabla \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) - \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \nabla \overline{\widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1)} \right) &= \\ &= -\delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_2) \overline{\widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1)} + \delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_1) \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \\ &= -\delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_2) \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{x}}_1)} + \delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_1) \widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \\ (\text{reciprocity}) &= -\delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_2) \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)} + \delta(\vec{\mathbf{y}} - \vec{\mathbf{x}}_1) \widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \end{split}$$

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The Kirchhoff Helmholtz identity

Proof continues

 $\bullet\,$ By integrating the previous expression over the ball B(0,L) and using the divergence theorem, we obtain

$$\begin{split} \int_{\partial B(0,L)} \frac{\vec{\mathbf{y}}}{|\vec{\mathbf{y}}|} \cdot \left(\widehat{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \nabla \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) - \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \nabla \overline{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \right) ds(\vec{\mathbf{y}}) \\ &= -\overline{\widehat{G}}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) + \widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \end{split}$$

• Using the Sommerfeld radiation condition (letting $L \to \infty$) we replace $\frac{\vec{\mathbf{y}}}{|\vec{\mathbf{y}}|} \cdot \nabla \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \text{ by } \frac{i\omega}{c_0} \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \text{ and}$ $\frac{\vec{\mathbf{y}}}{|\vec{\mathbf{y}}|} \cdot \nabla \overline{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \text{ by } \frac{-i\omega}{c_0} \overline{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \text{ to obtain}$

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The Kirchhoff Helmholtz identity

Proof continues

 $\bullet\,$ By integrating the previous expression over the ball B(0,L) and using the divergence theorem, we obtain

$$\begin{split} \int_{\partial B(0,L)} \frac{\vec{\mathbf{y}}}{|\vec{\mathbf{y}}|} \cdot \left(\widehat{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \nabla \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) - \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \nabla \overline{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \right) ds(\vec{\mathbf{y}}) \\ &= -\overline{\widehat{G}}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) + \widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \end{split}$$

• Using the Sommerfeld radiation condition (letting $L \to \infty$) we replace $\frac{\vec{\mathbf{y}}}{|\vec{\mathbf{y}}|} \cdot \nabla \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \text{ by } \frac{i\omega}{c_0} \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_2) \text{ and}$ $\frac{\vec{\mathbf{y}}}{|\vec{\mathbf{y}}|} \cdot \nabla \overline{\widehat{G}}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \text{ by } \frac{-i\omega}{c_0} \widehat{G}(\omega, \vec{\mathbf{y}}, \vec{\mathbf{x}}_1) \text{ to obtain}$ $\frac{2i\omega}{c_0} \int_{\partial B(0,L)} \overline{\widehat{G}}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}) \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) ds(\vec{\mathbf{y}}) = \widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) - \overline{\widehat{G}}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)$

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- $n(\vec{\mathbf{x}},t)$ is a zero mean stationary (in time) random process,
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- assume Gaussian statistics for the noise sources
- assume that the coherence time of the sources is small compared to the typical travel times, i.e., $F(t) = F_{\varepsilon}(t) = F\left(\frac{t}{\varepsilon}\right)$
- $K(\vec{\mathbf{x}})$ characterizes the spatial support of the sources (if $K(\vec{\mathbf{x}}) \equiv 1$ there are noise sources everywhere)

• Let $u(t, \vec{x}_1)$ and $u(t, \vec{x}_2)$ be the wave fields recorded at \vec{x}_1 and \vec{x}_2 . Their empirical cross correlation is

$$C_T(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \frac{1}{T} \int_0^T u(t, \vec{\mathbf{x}}_1) u(t+\tau, \vec{\mathbf{x}}_2) dt$$

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Theorem

The expectation of the empirical cross correlation C_T (with respect to the distribution of the sources) is independent of T:

$$\langle C_T(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \rangle = \frac{1}{2\pi} \int \int \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F_{\varepsilon}}(\omega) e^{-i\omega\tau} d\vec{\mathbf{y}} d\omega,$$

Proof

The solution of the wave equation can be written as (change of variable $\tilde{s} \rightarrow t - s$)

$$u(t, \vec{\mathbf{x}}) = \int_{-\infty}^{t} \int G(t-s, \vec{\mathbf{x}}, \vec{\mathbf{y}}) n_{\varepsilon}(s, \vec{\mathbf{y}}) d\vec{\mathbf{y}} ds = \int_{0}^{\infty} \int G(\tilde{s}, \vec{\mathbf{x}}, \vec{\mathbf{y}}) n_{\varepsilon}(t-\tilde{s}, \vec{\mathbf{y}}) d\vec{\mathbf{y}} d\tilde{s}$$

Extending $G(t, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = 0$ for $t \leq 0$, we get

$$u(t, \vec{\mathbf{x}}) = \int_{-\infty}^{\infty} \int G(s, \vec{\mathbf{x}}, \vec{\mathbf{y}}) n_{\varepsilon}(t - s, \vec{\mathbf{y}}) d\vec{\mathbf{y}} ds$$

The stationarity of n_{ε} implies stationarity of $u(t, \vec{\mathbf{x}})$, hence the mean of C_T does not depend on T and is given by

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) := \langle C_T(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \rangle = \langle u(0, \vec{\mathbf{x}}_1) u(\tau, \vec{\mathbf{x}}_2) \rangle$$

= $\int d\vec{\mathbf{y}}_1 \int d\vec{\mathbf{y}}_2 \int ds \int ds' G(s, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}_1) G(s', \vec{\mathbf{x}}_2, \vec{\mathbf{y}}_2) \langle n_{\varepsilon}(-s, \vec{\mathbf{y}}_1) n_{\varepsilon}(\tau - s', \vec{\mathbf{y}}_2) \rangle$
= $\int d\vec{\mathbf{y}} \int ds \int ds' G(s, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}) G(s', \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) F_{\varepsilon}(-s - \tau + s')$

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Proof

doing a change of variables $s_1 = s$, $s_2 = -s - \tau + s'$ we get

 $C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int d\vec{\mathbf{y}} \int ds_1 \int ds_2 G(s_1, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}) G(\tau + s_1 + s_2, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) F_{\varepsilon}(s_2)$

using Fourier transform and going in the frequency domain we obtain the result, $C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) =$

$$\begin{split} \int e^{-i\omega_1 s_1 - i\omega_2(\tau + s_1 + s_2)} \widehat{G}(\omega_1, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}) \widehat{G}(\omega_2, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) F_{\varepsilon}(s_2) d\vec{\mathbf{y}} ds_1 ds_2 \frac{d\omega_1}{2\pi} \frac{d\omega_2}{2\pi} \\ & (\int ds_1 \Rightarrow \omega_1 = -\omega_2) \\ \int e^{-i\omega_2(\tau + s_2)} \widehat{G}(-\omega_2, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}) \widehat{G}(\omega_2, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) F_{\varepsilon}(s_2) d\vec{\mathbf{y}} ds_2 \frac{d\omega_2}{2\pi} \\ & (\widehat{F}_{\varepsilon}(-\omega_2) = (\int e^{-i\omega_2 s_2} F_{\varepsilon}(s_2) ds_2) \text{ and } \widehat{F}_{\varepsilon} \text{ real valued and even} \\ & \int e^{-i\omega_2 \tau} \widehat{G}(-\omega_2, \vec{\mathbf{x}}_1, \vec{\mathbf{y}}) \widehat{G}(\omega_2, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F}_{\varepsilon}(\omega_2) d\vec{\mathbf{y}} \frac{d\omega_2}{2\pi} \\ & (f(t) \text{ real } \Rightarrow \widehat{f}(-\omega) = \overline{\widehat{f}(\omega)}) \\ & \int e^{-i\omega\tau} \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F}_{\varepsilon}(\omega) d\vec{\mathbf{y}} \frac{d\omega_2}{2\pi} \\ & = 1 \\ & =$$

We showed that the expectation of the empirical cross correlation C_T is independent of T and given by :

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \frac{1}{2\pi} \int \int \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F_{\varepsilon}}(\omega) e^{-i\omega\tau} d\vec{\mathbf{y}} d\omega,$$

We can re-write $C^{(1)}(au, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)$ as

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int e^{-i\omega\tau} \widehat{D}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \widehat{F}_{\varepsilon}(\omega) \frac{d\omega}{2\pi}$$

with

$$\widehat{D}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) d\vec{\mathbf{y}}$$

Theorem

The empirical cross correlation C_T is a self-averaging quantity

$$C_T(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \xrightarrow{T \to \infty} C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)$$

in probability with respect to the distribution of the sources.

- C_T is a statistically stable quantity : for large T, C_T is independent of the realization of the noise sources !
- This implies, in particular, that the SNR of the cross-correlations is proportional to $\sqrt{T}\,!$

$$SNR(X) = \frac{\langle X \rangle}{\sqrt{Var(X)}}$$

Proof: show that the variance of C_T tends to zero as $T \to \infty$ (the rate is O(1/T)).

The covariance of the empirical cross correlation C_T is :

$$Cov(C_T(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2), C_T(\tau', \vec{\mathbf{x}}_3, \vec{\mathbf{x}}_4)) = \frac{1}{2\pi T} \int \widehat{D}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_3) \overline{\widehat{D}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{x}}_4)} (\widehat{F_{\varepsilon}}(\omega))^2 e^{-i\omega(\tau'-\tau)} d\omega + \frac{1}{2\pi T} \int \widehat{D}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_4) \overline{\widehat{D}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{x}}_3)} (\widehat{F_{\varepsilon}}(\omega))^2 e^{-i\omega(\tau'+\tau)} d\omega$$

For details see

J. Garnier and G. Papanicolaou, "Passive sensor imaging using cross correlations of noisy signals in a scattering medium", SIAM J. Imaging Sciences, 2 :396–437, 2009.

• Assume that the medium is homogeneous outside B(0, D) and that the noise sources are distributed with uniform density on the surface of a sphere B(0, L) with $L \gg D$



- Assume that the medium is homogeneous outside B(0, D) and that the noise sources are distributed with uniform density on the surface of a sphere B(0, L) with $L \gg D$
- Then for any $\vec{\mathbf{x}}_1, \ \vec{\mathbf{x}}_2 \in B(0,D)$ we have

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int_{\partial B(0,L)} e^{-i\omega\tau} \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) \widehat{F}_{\varepsilon}(\omega) d\vec{\mathbf{y}} \frac{d\omega}{2\pi}$$

• Using the Kirchhoff-Helmholtz identity

 $\frac{2i\omega}{c_0}\int_{\partial B(0,L)}\overline{\widehat{G}(\omega,\vec{\mathbf{x}}_1,\vec{\mathbf{y}})}\widehat{G}(\omega,\vec{\mathbf{x}}_2,\vec{\mathbf{y}})ds(\vec{\mathbf{y}}) = \widehat{G}(\omega,\vec{\mathbf{x}}_1,\vec{\mathbf{x}}_2) - \overline{\widehat{G}(\omega,\vec{\mathbf{x}}_1,\vec{\mathbf{x}}_2)}$

we obtain (up to a multiplicative constant),

$$\frac{\partial}{\partial \tau} C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = -F_{\varepsilon} \star G(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) + F_{\varepsilon} \star G(-\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)$$

where \star denotes convolution in τ .

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• If $\varepsilon \ll 1$ then F_{ε} behaves like a delta-function and we get



• If $\varepsilon \ll 1$ then F_{ε} behaves like a delta-function and we get

 $\frac{\partial}{\partial \tau} C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) \approx G(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) - G(-\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2)$

• In the more general case, for spatially localised noise sources, the cross correlation between \vec{x}_1 and \vec{x}_2 is expected to have a singular component at the travel time between the two points only if the ray going through \vec{x}_1 and \vec{x}_2 reaches into the source region, that is, into the support of the function $K(\vec{y})$. This is shown using stationary phase.

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Spatially localised noise sources WKB assymptotics

 We assume a slowly varrying c(x) and use the WKB (Wentzell-Kramers-Brillouin) asymptotics for the Green's function

$$\widehat{G}(\frac{\omega}{\varepsilon}, \vec{\mathbf{x}}, \vec{\mathbf{y}}) \approx a(\vec{\mathbf{x}}, \vec{\mathbf{y}}) e^{i\omega \frac{\tau(\vec{\mathbf{x}}, \vec{\mathbf{y}})}{\varepsilon}}$$

The amplitude $a(\vec{\mathbf{x}}, \vec{\mathbf{y}})$ and travel time $\tau(\vec{\mathbf{x}}, \vec{\mathbf{y}})$ are smooth except at $\vec{\mathbf{x}} = \vec{\mathbf{y}}$.

• To obtain this we seek for an expansion of $\widehat{G}(\frac{\omega}{\varepsilon}, \vec{\mathbf{x}}, \vec{\mathbf{y}})$ as $\varepsilon \to 0$ of the form

$$\widehat{G}(\frac{\omega}{\varepsilon}, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = e^{i\frac{\omega}{\varepsilon}T(\vec{\mathbf{x}}, \vec{\mathbf{y}})} \sum_{j=0}^{\infty} A_j(\vec{\mathbf{x}}, \vec{\mathbf{y}}) \frac{\varepsilon^j}{\omega^j}$$

substituting in

$$\frac{\omega^2}{\varepsilon^2 c(\vec{\mathbf{x}})^2} \widehat{G}(\frac{\omega}{\varepsilon}, \vec{\mathbf{x}}, \vec{\mathbf{y}}) + \Delta \widehat{G}(\frac{\omega}{\varepsilon}, \vec{\mathbf{x}}, \vec{\mathbf{y}}) = -\delta(\vec{\mathbf{x}} - \vec{\mathbf{y}})$$

we get

Terms of
$$\mathcal{O}\left(\frac{1}{\varepsilon^2}\right)$$
: $|\nabla T(\vec{\mathbf{x}}, \vec{\mathbf{y}})|^2 - \frac{1}{c(\vec{\mathbf{x}})^2} = 0$, eikonal for the travel time

Terms of $\mathcal{O}\left(\frac{1}{\varepsilon}\right)$: $2\nabla T(\vec{\mathbf{x}}, \vec{\mathbf{y}}) \cdot \nabla A_0(\vec{\mathbf{x}}, \vec{\mathbf{y}}) + A_0(\vec{\mathbf{x}}, \vec{\mathbf{y}}) \Delta T(\vec{\mathbf{x}}, \vec{\mathbf{y}}) = 0$, transport for the amplitude

that we can solve with the method of characteristics (rays)

• We have $a(\vec{\mathbf{x}}, \vec{\mathbf{y}}) = A_0(\vec{\mathbf{x}}, \vec{\mathbf{y}})$ and $\tau(\vec{\mathbf{x}}, \vec{\mathbf{y}}) = T(\vec{\mathbf{x}}, \vec{\mathbf{y}})$

In the homogeneous case :
$$a(\vec{\mathbf{x}}, \vec{\mathbf{y}}) = \frac{1}{4\pi |\vec{\mathbf{x}} - \vec{\mathbf{y}}|}$$
 and $\tau(\vec{\mathbf{x}}, \vec{\mathbf{y}}) = \frac{|\vec{\mathbf{x}} - \vec{\mathbf{y}}|}{c_0}$

Ray equations

• The travel time can be obtained from Fermat's principle :

$$\tau(\vec{\mathbf{x}}, \vec{\mathbf{y}}) = \inf \left\{ Ts.t. \exists (X_t)_{t \in [0,T]} \in \mathcal{C}^1, X_0 = \vec{\mathbf{x}}, X_T = \vec{\mathbf{y}}, \left| \frac{dX_t}{dt} \right| = c(X_t) \right\}$$

• The minimising curve X_t is a ray and we assume that $c(\vec{\mathbf{x}})$ is such that there is a unique ray joining any $\vec{\mathbf{x}}, \vec{\mathbf{y}}$. We need the following Lemma :

Lemma

If $\nabla_{\vec{\mathbf{y}}}\tau(\vec{\mathbf{x}}_1,\vec{\mathbf{y}})=\nabla_{\vec{\mathbf{y}}}\tau(\vec{\mathbf{x}}_2,\vec{\mathbf{y}})$ then $\vec{\mathbf{x}}_1$ and $\vec{\mathbf{x}}_2$ lie on the same ray issuing from $\vec{\mathbf{y}}$ and

$$\tau(\vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = |\tau(\vec{\mathbf{x}}_2, \vec{\mathbf{y}}) - \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}})|.$$

While if $\nabla_{\vec{y}} \tau(\vec{x}_1, \vec{y}) = -\nabla_{\vec{y}} \tau(\vec{x}_2, \vec{y})$ then \vec{x}_1 and \vec{x}_2 lie on the opposite sides of the same ray issuing from \vec{y} and

$$\tau(\vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}}) + \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}}).$$

back to cross-correlations

• We have

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int e^{-i\omega\tau} \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F}_{\varepsilon}(\omega) d\vec{\mathbf{y}} \frac{d\omega}{2\pi}$$

• Now use that $\widehat{F}_{\varepsilon}(\omega) = \varepsilon \widehat{F}(\varepsilon \omega)$ (since $F_{\varepsilon}(t) = F(\frac{t}{\varepsilon})$) to get

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \varepsilon \int e^{-i\omega\tau} \overline{\widehat{G}(\omega, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\omega, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F}(\varepsilon\omega) d\vec{\mathbf{y}} \frac{d\omega}{2\pi}$$

• With the change of variables $\tilde{\omega} = \varepsilon \omega$,

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int e^{-i\frac{\tilde{\omega}}{\varepsilon}\tau} \overline{\widehat{G}(\frac{\tilde{\omega}}{\varepsilon}, \vec{\mathbf{x}}_1, \vec{\mathbf{y}})} \widehat{G}(\frac{\tilde{\omega}}{\varepsilon}, \vec{\mathbf{x}}_2, \vec{\mathbf{y}}) K(\vec{\mathbf{y}}) \widehat{F}(\tilde{\omega}) d\vec{\mathbf{y}} \frac{d\tilde{\omega}}{2\pi}$$

stationary phase

 $\bullet\,$ Using the WKB approximation for \widehat{G} we get

$$C^{(1)}(\tau, \vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \int \overline{a(\vec{\mathbf{x}}_1, \vec{\mathbf{y}})} a(\vec{\mathbf{x}}_2, \vec{\mathbf{y}}) e^{i\frac{\omega}{\varepsilon}T(\vec{\mathbf{y}})} K(\vec{\mathbf{y}}) d\vec{\mathbf{y}} \frac{d\omega}{2\pi}$$

with phase

$$\omega T(\vec{\mathbf{y}}) = \omega \left(\tau(\vec{\mathbf{x}}_2, \vec{\mathbf{y}}) - \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}}) - \tau \right)$$

• Stationary phase implies that the main contributions come from the critical points where

$$\frac{\partial}{\partial \omega} \left(\omega T(\vec{\mathbf{y}}) \right) = 0, \ \nabla_{\vec{\mathbf{y}}} \left(\omega T(\vec{\mathbf{y}}) \right) = 0$$

from where it follows that

$$\tau(\vec{\mathbf{x}}_2, \vec{\mathbf{y}}) - \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}}) = \tau, \ \nabla_{\vec{\mathbf{y}}} \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}}) = \nabla_{\vec{\mathbf{y}}} \tau(\vec{\mathbf{x}}_2, \vec{\mathbf{y}})$$

The previous Lemma implies that $\vec{\mathbf{x}}_1$ and $\vec{\mathbf{x}}_2$ lie on the same ray issuing from $\vec{\mathbf{y}}$! In order for a stationary point to contribute we need $K(\vec{\mathbf{y}}) \neq 0$ which means that $\vec{\mathbf{y}}$ has to be in the source region!

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Correlation based Imaging

Possible configurations when $\nabla_{\vec{y}} \tau(\vec{x}_1, \vec{y}) = \nabla_{\vec{y}} \tau(\vec{x}_2, \vec{y})$



 $\tau(\vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = \tau(\vec{\mathbf{x}}_2, \vec{\mathbf{y}}) - \tau(\vec{\mathbf{x}}_1, \vec{\mathbf{y}})$



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Introduction

- 2 Model problem
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- 8 Seasonal Variations
- 9 Experiments on Real data

Cross-correlations



Two seismic stations (Santorini and Naxos)

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Cross-correlations



The data recorded on Santorini filtered in frequency range [0.2, 0.5]Hz

Cross-correlations



Cross-correlations



The cross-correlations between the two stations as a function of recording time. Distance : 78Km , Mean Speed : 2.1 Km/s \Rightarrow expected peak at ± 37 s Recall SNR analysis (\sqrt{T})

Cross-correlations



The cross-correlations between the two stations as a function of recording time. Distance : 78Km , Mean Speed : 2.1 Km/s \Rightarrow expected peak at ± 37 s Recall SNR analysis (\sqrt{T})

Use these cross-correlations to : Estimate the velocity structure in the crust Volcano monitoring (Santorini/seismic activity) Seismic fault monitoring

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The imaging problem

• noise sources are spatially localized,

• sensors $(ec{\mathbf{x}}_q)_{1\leqslant q\leqslant N_{\mathrm{q}}}$ are located between the sources and the reflector



- This is the daylight configuration (Josselin's talk) : there are rays that emanating from the source region meet first a sensor and then the reflector !
- Stationary phase analysis shows that the cross correlation between two sensors \vec{x}_1 and \vec{x}_2 has a singular component at +/- the sum of the travel times between the sensors and the reflector

$$\tau(z^*, \vec{\mathbf{x}}_1) + \tau(z^*, \vec{\mathbf{x}}_2)$$

Imaging functional

• To image the reflector it seems a good idea to compute

$$\mathcal{I}^{\mathrm{D}}(\mathbf{z}) = \sum_{q,q'=1}^{N_{\mathrm{q}}} C_T(\tau(\mathbf{z}, \vec{\mathbf{x}}_{q'}) + \tau(\mathbf{z}, \vec{\mathbf{x}}_{q}), \vec{\mathbf{x}}_{q}, \vec{\mathbf{x}}_{q'}),$$

for points \mathbf{z} in a search domain.

- we could also use the $-(\tau(\mathbf{z}, \vec{\mathbf{x}}_{q'}) + \tau(\mathbf{z}, \vec{\mathbf{x}}_{q})))$, but it should not make a difference since we expect $C_T(-\tau, \vec{\mathbf{x}}_q, \vec{\mathbf{x}}_{q'}) = C_T(\tau, \vec{\mathbf{x}}_{q'}, \vec{\mathbf{x}}_q)$. In practice we may want to average the cross-correlation over the positive and negative times.
- $\mathcal{I}^{D}(\mathbf{z})$ is Kirchhoff migration (or travel time migration) and has been analyzed in detail (Beylkin, Bleistein, Symes, ...) in the case of an active array, i.e., when we send pulses from sources at the array and record the echoes at the receivers.
- Here the difference is that the array is passive and records noisy signals. By forming the cross-correlations of the recorded signals we transform the "passive array" into an active one.

Simulation setup

- wave equation on the rectangle $[0, 50\lambda] \times [-15\lambda, 15\lambda]$, with a reflector located on $[44\lambda, 46\lambda] \times [-\lambda, \lambda]$,
- random distribution of sources has support on the rectangle $[0, 4\lambda] \times [-15\lambda, 15\lambda]$,
- we record the solution u of the wave equation at $N_{\rm q}$ receivers located at

$$\label{eq:constraint} \begin{split} \vec{\mathbf{x}}_q &= (5\lambda, (q-(N_{\rm q}+1)/2)\lambda/2)\text{,}\\ \text{for } 1 \leqslant q \leqslant N_{\rm q}\text{,} \end{split}$$

- $\lambda = 6$ mm and $c_0 = 3$ km/s,
- the reflector is modeled as a soft acoustic scatterer, *i.e.* u = 0 on the boundary of the reflector
- we surround domain by PML

4 6 6	a=30λ	L=39 λ	
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Image domain : square of size 20λ centered on the reflector

• We solve the wave equation with the code Montjoie (http://montjoie.gforge.inria.fr/).

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- Montjoie is designed for the efficient solution of time-domain and time-harmonic linear partial differential equations using high-order finite element methods.

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- We solve the wave equation with the code Montjoie (http://montjoie.gforge.inria.fr/).
- Montjoie is designed for the efficient solution of time-domain and time-harmonic linear partial differential equations using high-order finite element methods.
- For the numerical examples considered here we use 7th order finite elements in space and 4th order finite differences in time.
- We added the computation of cross-correlations and imaging functionals in Montjoie.

Results (homogeneous test case)



Figure : Imaging functional for the homogeneous medium. $N_{\rm q}=21$

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Figure : Imaging functional for the homogeneous medium. $N_{\rm q}=31$

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Results (homogeneous test case)



Figure : Imaging functional for the homogeneous medium. $N_{\rm q}=41$

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Figure : Imaging functional for the homogeneous medium. $N_{\rm q}=51$

C. Tsogka (University of Crete)
• A good question that naturally arise is "Under what conditions do we obtain such a good image" ?

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- A good question that naturally arise is "Under what conditions do we obtain such a good image" ?
- In other words, "What are the parameters that control the quality of the image, and how" ?

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- A good question that naturally arise is "Under what conditions do we obtain such a good image" ?
- In other words, "What are the parameters that control the quality of the image, and how" ?
- Resolution analysis (same as for the active case)
 - Cross range resolution : $\lambda L/a$.
 - Range resolution : c_0/B

J. Garnier and G. Papanicolaou, "Resolution analysis for imaging with noise", Inverse Problems, Vol. 26, pp. 074001, 2010.

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• SNR analysis : Assuming that the sources surround the region of interest, and that the receiver array is sampled at half-a-wavelength apart (or larger), we show that

$$\mathrm{SNR}^{\mathrm{D}} = \frac{\left\langle \mathcal{I}^{\mathrm{D}}(\mathbf{z}_{\mathrm{r}}) \right\rangle}{\mathrm{Var} \left(\mathcal{I}^{\mathrm{D}}(\mathbf{z}) \right)^{1/2}} \sim \frac{N_{\mathrm{q}}^2 B}{\sqrt{N_{\mathrm{q}}^2 B/T}} = N_{\mathrm{q}} \sqrt{BT}$$

• Numerically we compute

$$\mathsf{SNR} = \frac{|\mathcal{I}^{\mathsf{D}}|(\mathbf{z}^*)}{\max_{\mathbf{z}\neq\mathbf{z}^*}|\mathcal{I}^{\mathsf{D}}|(\mathbf{z})}$$

where z^* is the point where the image admits its maximal value and $z \neq z^*$ means that squares of size $2\lambda \times 2\lambda$ centered at z and z^* do not intersect.

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Array size



Resolution improves with array size and SNR increases linearly with number of receivers

Recording time



SNR is proportional to \sqrt{T}

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Bandwidth



Range resolution improves with bandwidth and SNR is proportional to \sqrt{B}

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Our analysis shows that the important parameters for imaging are :

- The number of sensors $N_{\rm q}$. Both cross-range resolution and SNR linearly improve with $N_{\rm q}$.
- **②** The bandwidth of the noise sources B. Range resolution improves linearly with B, while the SNR is proportional to \sqrt{B} .
- The recording time T. The SNR of the cross-correlations, and therefore the SNR of the image as well, is proportional to \sqrt{T} .

Numerical results are in very good agreement with the theory.

J. Garnier, G. Papanicolaou, A. Semin and C.T., "Signal to Noise Ratio estimation in passive correlation based imaging", SIAM J. Imaging Sci. 6-2 (2013), pp. 1092-1110.

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 motivation : exploration geophysics

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- motivation : exploration geophysics
- traditional approach : surface array imaging

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- motivation : exploration geophysics
- traditional approach : surface array imaging
- complex medium impedes the imaging process

problem : small signal to noise ratio, i.e., scatterer echoes are overwhelmed by reflections from the background medium.



- motivation : exploration geophysics
- traditional approach : surface array imaging
- complex medium impedes the imaging process
- auxiliary receiver array does it helps and how ?

How to use the recorded signals on the auxiliary array so as to make a good image?



- motivation : exploration geophysics
- traditional approach : surface array imaging
- complex medium impedes the imaging process
- auxiliary receiver array does it helps and how ?

Mathematical Model : study wave propagation in inhomogeneous random media with a velocity that fluctuates arround a known mean

- Here there are no random sources, the sources are deterministic
- the randomness comes form the complex medium. Here $\mu_{\rm r}$ is obtained by combining an isotropic and a layered random variable,

$$\mu(\vec{\mathbf{x}}) = \frac{1}{\sqrt{2}} \left(\mu_i(\vec{\mathbf{x}}) + \mu_l(\vec{\mathbf{x}}) \right),$$

with standard deviation $\sigma = 0.08$. The isotropic part $\mu_i(\vec{\mathbf{x}}) = \mu\left(\frac{x}{\ell}, \frac{z}{\ell}\right)$, has a Gaussian correlation function

$$E\{\mu_i(\vec{\mathbf{x}}_1)\mu_i(\vec{\mathbf{x}}_2)\} = R(\vec{\mathbf{x}}_1, \vec{\mathbf{x}}_2) = e^{-\frac{|\vec{\mathbf{x}}_1 - \vec{\mathbf{x}}_2|^2}{2\ell^2}}, \qquad \ell = \lambda/2$$

and the layered random variable, $\mu_l(\vec{\mathbf{x}}) = \mu\left(\frac{z}{\ell_z}\right)$, satisfies

$$E\{\mu_l(z_1)\mu_l(z_2)\} = \left(1 + \frac{|z_1 - z_2|}{\ell_z}\right) e^{-\frac{|z_1 - z_2|}{\ell_z}}, \quad \ell_z = \lambda/30.$$

Image: A mathematical states and a mathem

Traditional Imaging



Statistically unstable and with very low signal to noise ratio : good image only for very low fluctuations

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Traditional Imaging



Statistically unstable and with very low signal to noise ratio : good image only for very low fluctuations

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• We compute cross-correlations at the auxiliary array by

$$C(\tau, \vec{\mathbf{x}}_l, \vec{\mathbf{x}}_m) = \frac{1}{T} \sum_{s=1}^{N_s} \int u(t, \vec{\mathbf{x}}_l; \vec{\mathbf{x}}_s) u(t+\tau, \vec{\mathbf{x}}_m; \vec{\mathbf{x}}_s) dt$$

• Note that this is coherent averaging over the sources which is not what we do in the ambient noise problem where

$$C(\tau, \vec{\mathbf{x}}_l, \vec{\mathbf{x}}_m) = \frac{1}{T} \int \left(\sum_{s=1}^{N_{\rm s}} u(t, \vec{\mathbf{x}}_l; \vec{\mathbf{x}}_s) \right) \left(\sum_{s=1}^{N_{\rm s}} u(t+\tau, \vec{\mathbf{x}}_m; \vec{\mathbf{x}}_s) \right) dt$$

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Correlation based Imaging



Correlation based imaging with the auxiliary array

Statistically stable : high SNR even for strong medium fluctuations (depends on bandwidth, complex medium characteristics, array characteristics)

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Wave propagation through a strongly scattering medium

• The field transmitted through the random strongly scattering medium has Gaussian statistics, mean zero, and cross correlations of the form

$$\begin{split} \mathbb{E} \Big[\widehat{u}(\omega, \vec{\mathbf{x}}_q; \vec{\mathbf{x}}_s) \widehat{u}(\omega', \vec{\mathbf{x}}_{q'}; \vec{\mathbf{x}}_{s'}) \Big] &= 0 \\ \mathbb{E} \Big[\widehat{u}(\omega, \vec{\mathbf{x}}_q; \vec{\mathbf{x}}_s) \overline{\widehat{u}(\omega', \vec{\mathbf{x}}_{q'}; \vec{\mathbf{x}}_{s'})} \Big] &= \widehat{F}(\omega) \overline{\widehat{F}(\omega')} \exp\left(- \frac{(\omega - \omega')^2}{\omega_c^2} \right) \\ &\times \exp\left(- \frac{|\mathbf{x}_q - \mathbf{x}_{q'}|^2}{X_{cq}^2} - \frac{|\mathbf{x}_s - \mathbf{x}_{s'}|^2}{X_{cs}^2} \right), \end{split}$$

- $\omega_{\rm c}$: correlation frequency of the field in the plane z=-L of the auxiliary receiver array,
- X_{cq} is the correlation radius of the field at the auxiliary receiver array (when emitted from a point source at the source array),
- X_{cs} is the correlation radius of the field at the source array (when emitted from a point source at the auxiliary receiver array).

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Image quality : Resolution analysis

Cross range resolution : $\lambda(L_y - L)/a_{\text{eff}}$.

with
$$a_{ ext{eff}}^{ ext{hom}} = b rac{L_y - L}{L}$$
 and $a_{ ext{eff}}^r = rac{\lambda(L_y - L)}{X_{ ext{cq}}}$

and is improved in the random paraxial case.



(a) Homogeneous medium

(b) Random medium

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J. Garnier and G. Papanicolaou, Role of scattering in virtual source imaging, SIAM Journal of Imaging Science 7 (2014), pp. 1210-1236.

C. Tsogka (University of Crete)

Cross-range resolution : $\lambda (L_y - L)/a_{ ext{eff}}$



The array size does not seem to affect the resolution! (here $a_{eff} = 9\lambda$)

Range resolution : c_0/B



Range resolution improves with B

We show that, the SNR defined by,

$$SNR_{CC} = \frac{|\mathbb{E}[\mathcal{I}_{CC}(\vec{\mathbf{y}})]|}{Var(\mathcal{I}_{CC}(\vec{\mathbf{y}}^S))^{1/2}},$$
(1)

satisfies (for $X_{\rm cs} \ll b$ and $\omega_c \ll B$)

$$\mathrm{SNR}_{\mathrm{CC}} \approx \frac{\sigma_{\mathrm{ref}} X_{\mathrm{cq}}}{\lambda^2 (L_y - L)} \Big(\frac{b}{\max\{\Delta X_{\mathrm{s}}, X_{\mathrm{cs}}\}} \Big)^{1/2} \Big(\frac{a}{\max\{\Delta X_{\mathrm{q}}, X_{\mathrm{cq}}\}} \Big) \Big(\frac{B}{\omega_{\mathrm{c}}} \Big)^{1/2}.$$

When the correlation radius X_{cq} is small, it is relevant to have a dense auxiliary receiver array for a given aperture in order to get good stability.

For a given aperture a, the SNR increases when the inter-distance $\Delta X_{\rm q}$ decreases, until the inter-distance becomes of the order of the correlation radius $X_{\rm cq}$ of the illumination field, and then the SNR reaches a value determined by $X_{\rm cq}$.

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SNR vs bandwidth



Figure : SNR vs bandwidth

• SNR is proportional to \sqrt{B}

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SNR vs number of sources



Figure : SNR as a function of b (in λ). Blue circles correspond to $\Delta X_{\rm s} = \lambda/2$, red squares to $\Delta X_{\rm s} = \lambda$ and yellow triangles to $\Delta X_{\rm s} = 3\lambda/2$.

• SNR is proportional to \sqrt{b} ($X_{\rm cs} > \Delta X_{\rm s}$)

Image: A mathematical states and a mathem

SNR vs number of receivers



• SNR is proportional to $N_{
m q}~(X_{
m cq} < \Delta X_{
m q})$

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Virtual source array imaging : Imaging by migrating the cross-correlation matrix of the auxiliary array data gives much better images then migrating the data !

In fact the random medium effects are eliminated and it is "even better" then if an active array near the object was used for imaging (only the down-going part of the Green's function is reconstructed)

Differences w.r.t the homogeneous case :

- SNR depends on the inter-element (auxiliary) array distance which should be of the order of X_{cq} for optimal SNR.
- SNR is proportional to $\sqrt{N_s}$. SNR also depends on the inter-element (source) array distance which should be of the order of X_{cs} for optimal SNR.

Numerical results are in very good agreement with the theory. Resolution analysis in agreement with results obtained in the random paraxial regime.

J. Garnier, G. Papanicolaou, A. Semin and C.T., Signal to Noise Ratio estimation in virtual source array imaging, SIAM Journal on Imaging Science 8 (2015), pp. 248-279.

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The seismic stations



Two seismic stations (Santorini and Naxos)

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Recordings on Santorini station



The data recorded on Santorini filtered in frequency range [0.2, 0.5]Hz

Recordings on Naxos station



The data recorded on Naxos filtered in frequency range [0.2, 0.5]Hz

Cross-correlations (3 months)



Observe variability w.r.t to the season of the year ! Due to seasonal variations of the medium or of the noise sources ? (non-stationarity)

C. Tsogka (University of Crete)

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We want to use the cross-correlations to estimate relative velocity changes in the propagating medium.

Two methods are mainly used to estimate dc/c. They both compare :

 $\mathrm{CC}_{\mathrm{ref}}$: the average CC over a long period (reference)

 $\mathrm{CC}_{\mathrm{cur}}$: average of CC over a small period around the current day

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Two methods are mainly used to estimate dc/c. They both compare :

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- MWCS method operates in the frequency domain
- SM operates in the time domain

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A small change in the velocity $c \to c(1 + dc/c)$ induces a change in travel time $t \to t(1 + dt/t)$ (with dc/c = -dt/t). We want to measure the stretching coefficient $\varepsilon = dt/t$ by comparing CC_{ref} and CC_{cur} .

We seek ε that maximizes :

$$C(\varepsilon) = \frac{\int \mathrm{CC}_{\mathrm{cur}}(t(1+\varepsilon))\mathrm{CC}_{\mathrm{ref}}(t)dt}{\sqrt{(\int \mathrm{CC}_{\mathrm{cur}}^2(t(1+\varepsilon))dt)(\int \mathrm{CC}_{\mathrm{ref}}^2(t)dt)}}$$

where the integration is over a specific time window.

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where the integration is over a specific time window. We seek for the parameter ε in a search interval $[-\varepsilon_{max} : \varepsilon_{max}]$, with an accuracy $d\varepsilon = 10^{-6}$ (using adaptive refinement).

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The MWCS method detects a phase shift in the frequency domain.

We divide each cross-correlation function into N_w windows, with each window centered around time t_i , $i = 1, ..., N_w$. For each central time t_i we get a measurement dt_i using the corresponding windowed segments of CC_{ref} and CC_{cur} .



Two methods to estimate $dc/c_{\rm MWCS\ Method}$



Those segments after being window tapered, they are being fourier transformed and called $F_{\rm ref}(\nu) = \mathcal{F}(\mathrm{CC}_{\rm ref})$ and $F_{\rm cur}(\nu) = \mathcal{F}(\mathrm{CC}_{\rm ref})$ respectively. Then the cross-spectrum is calculated as

$$X(\nu) = F_{\rm ref}(\nu) \ F^*_{\rm cur}(\nu),$$

Two methods to estimate $dc/c_{\rm MWCS\ Method}$

We estimate the phase of the cross-spectrum $\varphi_i(\nu_j)$ as a function of frequency ν_j ,

 $\varphi_i(\nu_j) = 2\pi \ dt_i \ \nu_j,$



and then we estimate dt_i the time shift corresponding to the central time t_i . From this we get dc/c = -dt/t using weighted least squares.

D. Clarke and L. Zaccarelli and N.M. Shapiro and F. Brenguier "Assessment of resolution and accuracy of the Moving Window Cross Spectral technique for monitoring crustal temporal variations using ambient seismic noise", Geophys. J. Int. 186 (2011) 867-882.

C. Tsogka (University of Crete)

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What are seasonal variations?

It has been observed that noisy recordings have seasonal variations



Power spectral density of recordings on Milos island. Vertical axis : time in days (850 days ≈ 2.3 years). Horizontal axis : frequency between 0 and 1 Hz.

Seasonal variations in the seismic velocity are observed in

Ueli Meier, Nikolai M. Shapiro and Florent Brenguier, "Detecting seasonal variations in seismic velocities within Los Angeles basin from correlations of ambient seismic noise", Geophys. J. Int. (2010) 181, 985–996

where it is suggested that they are due to hydrological and/or thermoelastic variations of the medium.

On the other hand

Zhongwen Zhan, Victor C. Tsai and Robert W. Clayton, "Spurious velocity changes caused by temporal variations in ambient noise frequency content", Geophys. J. Int. (2013) 194, 1574–1581

suggest that the observed variations are spurious and are due to seasonal variations in the amplitude spectrum of the noise sources.

Seasonal variations in the seismic velocity are observed in

Ueli Meier, Nikolai M. Shapiro and Florent Brenguier, "Detecting seasonal variations in seismic velocities within Los Angeles basin from correlations of ambient seismic noise", Geophys. J. Int. (2010) 181, 985–996

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To study the effect of seasonal variations of the ambient noise sources on the estimated changes of the seismic velocity we design simple but realistic numerical experiments

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Modelling the seasonal variations : the wave equation

We consider the acoustic wave equation :

$$\frac{1}{c(\vec{\mathbf{x}})^2}\frac{\partial^2 u}{\partial t^2}(t,\vec{\mathbf{x}}) - \Delta u(t,\vec{\mathbf{x}}) = n(t,\vec{\mathbf{x}}),$$

where $n(t, \vec{\mathbf{x}})$ models the noise sources.

Image: A matrix

We consider the acoustic wave equation :

$$\frac{1}{c(\vec{\mathbf{x}})^2}\frac{\partial^2 u}{\partial t^2}(t,\vec{\mathbf{x}}) - \Delta u(t,\vec{\mathbf{x}}) = n(t,\vec{\mathbf{x}}),$$

where $n(t, \vec{\mathbf{x}})$ models the noise sources. The solution at a given point $\vec{\mathbf{x}}$ can be written as,

$$u(t, \vec{\mathbf{x}}) = \int \int G(t - s, \vec{\mathbf{x}}, \vec{\mathbf{y}}) n(s, \vec{\mathbf{y}}) d\vec{\mathbf{y}} ds,$$

or equivalently in the frequency domain,

$$\hat{u}(\omega, \vec{\mathbf{x}}) = \int \hat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}) \hat{n}(\omega, \vec{\mathbf{y}}) d\vec{\mathbf{y}}.$$

Here hat denotes the Fourier transform and $\hat{G}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}})$ is the Green's function.

We assume that $n(t,\vec{\mathbf{x}})$ is a zero-mean stationary in time random process with a covariance function

$$\langle n(t_1, \vec{\mathbf{y}}_1), n(t_2, \vec{\mathbf{y}}_2) \rangle = \Gamma(t_2 - t_1, \vec{\mathbf{y}}_1) \delta(\vec{\mathbf{y}}_2 - \vec{\mathbf{y}}_1).$$

Here $\langle \cdot \rangle$ stands for statistical averaging. The function $t \to \Gamma(t, \vec{y})$ is the time correlation function of the noise signals emitted by the noise sources at location \vec{y} .

We assume that $n(t,\vec{\mathbf{x}})$ is a zero-mean stationary in time random process with a covariance function

$$\langle n(t_1, \vec{\mathbf{y}}_1), n(t_2, \vec{\mathbf{y}}_2) \rangle = \Gamma(t_2 - t_1, \vec{\mathbf{y}}_1) \delta(\vec{\mathbf{y}}_2 - \vec{\mathbf{y}}_1).$$

Here $\langle \cdot \rangle$ stands for statistical averaging. The function $t \to \Gamma(t, \vec{\mathbf{y}})$ is the time correlation function of the noise signals emitted by the noise sources at location $\vec{\mathbf{y}}$.

The function $\vec{\mathbf{y}} \to \Gamma(0, \vec{\mathbf{y}})$ characterizes the spatial support of the sources. In our case we assume that the sources are uniformly distributed on a circle C of radius of $R_C = 25 \text{km}$:

$$\Gamma(t, \vec{\mathbf{y}}) = \frac{1}{2\pi R_{\mathcal{C}}} \Gamma_0(t, \vec{\mathbf{y}}) \delta_{\mathcal{C}}(\mathbf{y}).$$

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Modelling the seasonal variations : setup

The noise sources are located on a circle of radius 25km and the wave field is recorded at two receivers $\vec{x}_1 = (-5, 0)$ km and $\vec{x}_2 = (5, 0)$ km.



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To model seasonal variations we assume that the statistics of the noise sources change from one day to another and denote $\Gamma_0^j(t, \vec{\mathbf{y}})$ the covariance function at day j. The wave field at $\vec{\mathbf{x}}$ is computed by $(i = 1, N_s \text{ the noise sources})$

$$\hat{u}^{j}(\omega, \vec{\mathbf{x}}) = \frac{1}{N_{s}} \sum_{i=1}^{N_{s}} \hat{G}^{j}(\omega, \vec{\mathbf{x}}, \vec{\mathbf{y}}_{i}) \hat{n}_{i}^{j}(\omega),$$
(2)

where $\hat{n}_i^j(\omega)$ is the frequency content of the noise sources at $\vec{\mathbf{y}}_i$ during day j, which is random such that $\left<\hat{n}_i^j(\omega)\right>=0$ and

$$\left\langle \hat{n}_{i}^{j}(\omega)\overline{\hat{n}_{i}^{j}}(\omega')\right\rangle = 2\pi\hat{\Gamma}_{0}^{j}(\omega,\vec{\mathbf{y}}_{i})\delta(\omega-\omega').$$

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Our model for the power spectral density of the noise sources is

$$\hat{\Gamma}^{j}_{0}(\omega, \vec{\mathbf{y}}) = \hat{F}(\omega)\hat{s}^{j}(\omega, \vec{\mathbf{y}}),$$

Here the unperturbed noise source distribution is uniform over the circle C and has power spectral density $\hat{F}(\omega)$, and $\hat{s}^{j}(\omega, \vec{\mathbf{y}})$ is the daily perturbation of the power spectral density at location $\vec{\mathbf{y}}$.

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For uniform daily perturbations, we have $\hat{s}^{j}(\omega, \vec{\mathbf{y}}) = \hat{f}^{j}(\omega)l(\vec{\mathbf{y}}),$

$$\left\langle \hat{CC}^{j}(\omega, \vec{\mathbf{x}}_{1}, \vec{\mathbf{x}}_{2}) \right\rangle = \hat{F}(\omega) \hat{f}^{j}(\omega) \int_{\mathcal{C}} d\sigma(\vec{\mathbf{y}}) \ \overline{\hat{G}^{j}(\omega, \vec{\mathbf{x}}_{1}, \vec{\mathbf{y}})} \hat{G}^{j}(\omega, \vec{\mathbf{x}}_{2}, \vec{\mathbf{y}}) l(\vec{\mathbf{y}}),$$

and the daily perturbation factors out of the integral ...

Results

Reference CC-function = average 360 daily CC-functions Current CC-function for day j = average 7 days around day j.

Without seasonal variations



Results



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Results

With seasonal variations



We clearly see that only the SM estimation is affected by the seasonal variations in the amplitude spectra of the noise sources.

C. Tsogka (University of Crete)

Correlation based Imaging

Spectral Whitening

A simple way to remove them is to use spectral whitening (normalize the amplitude spectra) of the \mathbb{C} -functions.



Spectral whitening works for perturbations of separable form, i.e., uniform seasonal variations (a reasonable assumption for receivers that are not very far apart)

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- Estimating velocity change in a medium
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- Experiments on Real data



We consider Milos, an island in Aegean sea and two stations 6 Km apart.

Red : dc/c estimation without using Spectral Whitening. Blue : dc/c estimation using Spectral Whitening

SNR improvement of an order of 3 (std).

Santorini 2011-2012 unrest

The Santorini 2011-2012 seismic unrest begins on January 2011 and ends on February 2012 measuring a total of 10 cm uplift of the caldera of Santorini.



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Santorini 2011-2012 unrest : dc/c estimation

Slow event, difficult to follow without removing seasonal variations !

We observe decrease in the velocity of seismic waves in the caldera of Santorini which is correlated with the accumulated elevation measured with GPS.



Conclusions/Future work

- Potential of developing monitoring tools which provide accurate results even with sparse seismic networks
 - E. Daskalakis, C. Evangelidis, J. Garnier, N. Melis, G. Papanicolaou and C. T., "Robust seismic velocity change estimation using ambient noise recordings", preprint 2015.
- Study how errors propagate into the estimation (UQ)
- Application to seismic fault monitoring

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Future work

Results obtained with the SM method when in the computation of $\mathrm{CC}_{\it cur}$ we only use days in the past.



On the left we use a window of 15 days and on the right 21 days. The day of the earthquake (M6.9) is shown with the blue vertical curve. We clearly observe a decrease in the relative velocity of about 0.4% starting a few days before and continuing for several days after the earthquake. Similar results have been obtained for the Tohuku-Oki earthquake in Japan

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